TU Delft

Delft University of Technology

Document Version
Final published version

Licence
CC BY

Citation (APA)
Xu, Y., Wang, X., Ou, Q., Zhou, Z., van der Hoek, J. P., & Liu, G. (2024). Appearance of Recalcitrant Dissolved Black Carbon and Dissolved Organic Sulfur in River Waters Following Wildfire Events. *Environmental Science and Technology*, *58*(16), 7165-7175. https://doi.org/10.1021/acs.est.4c00492

Important note
To cite this publication, please use the final published version (if applicable).
Please check the document version above.

Copyright
In case the licence states “Dutch Copyright Act (Article 25fa)”, this publication was made available Green Open Access via the TU Delft Institutional Repository pursuant to Dutch Copyright Act (Article 25fa, the Taverne amendment). This provision does not affect copyright ownership.
Unless copyright is transferred by contract or statute, it remains with the copyright holder.

Sharing and reuse
Other than for strictly personal use, it is not permitted to download, forward or distribute the text or part of it, without the consent of the author(s) and/or copyright holder(s), unless the work is under an open content license such as Creative Commons.

Takedown policy
Please contact us and provide details if you believe this document breaches copyrights.
We will remove access to the work immediately and investigate your claim.

This work is downloaded from Delft University of Technology.

Environmental Science & Technology

pubs.acs.org/est

Open Access

This article is licensed under CC-BY 4.0

Article

Appearance of Recalcitrant Dissolved Black Carbon and Dissolved Organic Sulfur in River Waters Following Wildfire Events

Published as part of Environmental Science & Technology virtual special issue "Wildland Fires: Emissions, Chemistry, Contamination, Climate, and Human Health".

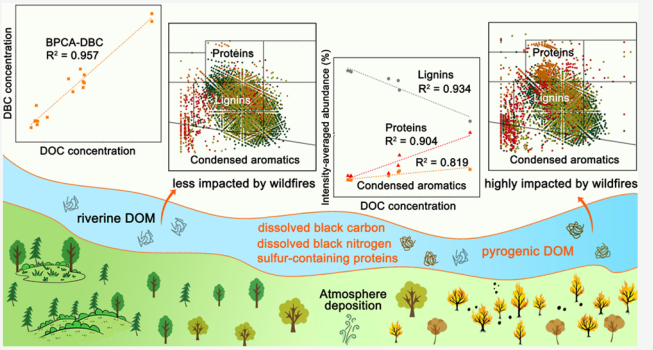
Yanghui Xu,[▽] Xintu Wang,[▽] Qin Ou, Zhongbo Zhou,* Jan Peter van der Hoek, and Gang Liu*

Cite This: https://doi.org/10.1021/acs.est.4c00492

Read Online

ACCESS | Metrics & More | Article Recommendations | Supporting Information

ABSTRACT: Increasing wildfire frequency, a consequence of global climate change, releases incomplete combustion byproducts such as aquatic pyrogenic dissolved organic matter (DOM) and black carbon (DBC) into waters, posing a threat to water security. In August 2022, a series of severe wildfires occurred in Chongqing, China. Samples from seven locations along the Yangtze and Jialing Rivers revealed DBC, quantified by the benzene poly(carboxylic acid) (BPCA) method, comprising 9.5−19.2% of dissolved organic carbon (DOC). High concentrations of BPCA-DBC with significant polycondensation were detected near wildfire areas, likely due to atmospheric deposition driven by wind. Furthermore, Fourier transform ion cyclotron resonance mass spectrometry (FT-ICR-MS) revealed that wildfires were associated with an increase in condensed aromatics, proteins, and unsaturated hydrocarbons, along with a decrease in lignins. The condensed aromatics primarily consisted of dissolved black nitrogen (DBN), contributing to abundant high-nitrogen-containing compounds in locations highly affected by wildfires. Meanwhile, wildfires potentially induced the input of recalcitrant sulfur-containing protein-like compounds, characterized by high oxidation, aliphatic nature, saturation, and low aromaticity. Overall, this study revealed the appearance of recalcitrant DBC and dissolved organic sulfur in river waters following wildfire events, offering novel insights into the potential impacts of wildfires on water quality and environmental biogeochemistry.



KEYWORDS: *wildfires, dissolved organic matter, dissolved black carbon, dissolved black nitrogen, sulfur-containing proteins*

Downloaded via TU DELFT on April 19, 2024 at 09:17:47 (UTC).
See https://pubs.acs.org/sharingguidelines for options on how to legitimately share published articles.

INTRODUCTION

The rising frequency and severity of wildfires, closely tied to climate change, are expected to further escalate as climate change persists.[1,2] Wildfires emit significant amounts of carbon into the atmosphere in the form of CO_2 due to the combustion of biomass.[1] Wildfires also produce substantial quantities of pyrogenic organic matter (OM) during the incomplete combustion of vegetation and soil organic material.[3,4] Rivers originating in fire-affected areas can undergo significant alterations during and after burn events, impacting water quality and freshwater ecosystem productivity.[5,6] Particularly, the export of pyrogenic OM in the form of dissolved organic matter (DOM) into rivers has the potential to adversely affect water quality and can influence carbon biogeochemical cycling and aquatic ecosystem functioning.[7,8] Dissolved black carbon (DBC), the largest known refractory DOM pool in aquatic systems, has arisen great attention for its roles in regulating greenhouse gas emissions and stabilizing DOM.[9,10]

The relationship between wildfires and riverine DOM/DBC has been a research hotspot, but it remains not well understood. Ding et al. reported that there was no correlation between recent fire activities and in-stream dissolved organic carbon (DOC) and DBC concentrations.[11,12] However, Dittmar et al. observed continued export of riverine DBC and DOC each year in the rainy season after major fire activity.[7] The concentrations of DOC in streamwater were also identified to be primarily driven by local hydrology during the first storm event occurring 1 month after the fire.[13] Generally, most studies have investigated the correlation between wildfires and riverine DOM/DBC over an extended period, occurring several months or years after the wildfires took place.[7,14] Environmental drivers, especially hydrology during rain and storm events are recognized as significant factors

Received: January 14, 2024
Revised: March 29, 2024
Accepted: April 1, 2024
Published: April 10, 2024

© 2024 The Authors. Published by American Chemical Society

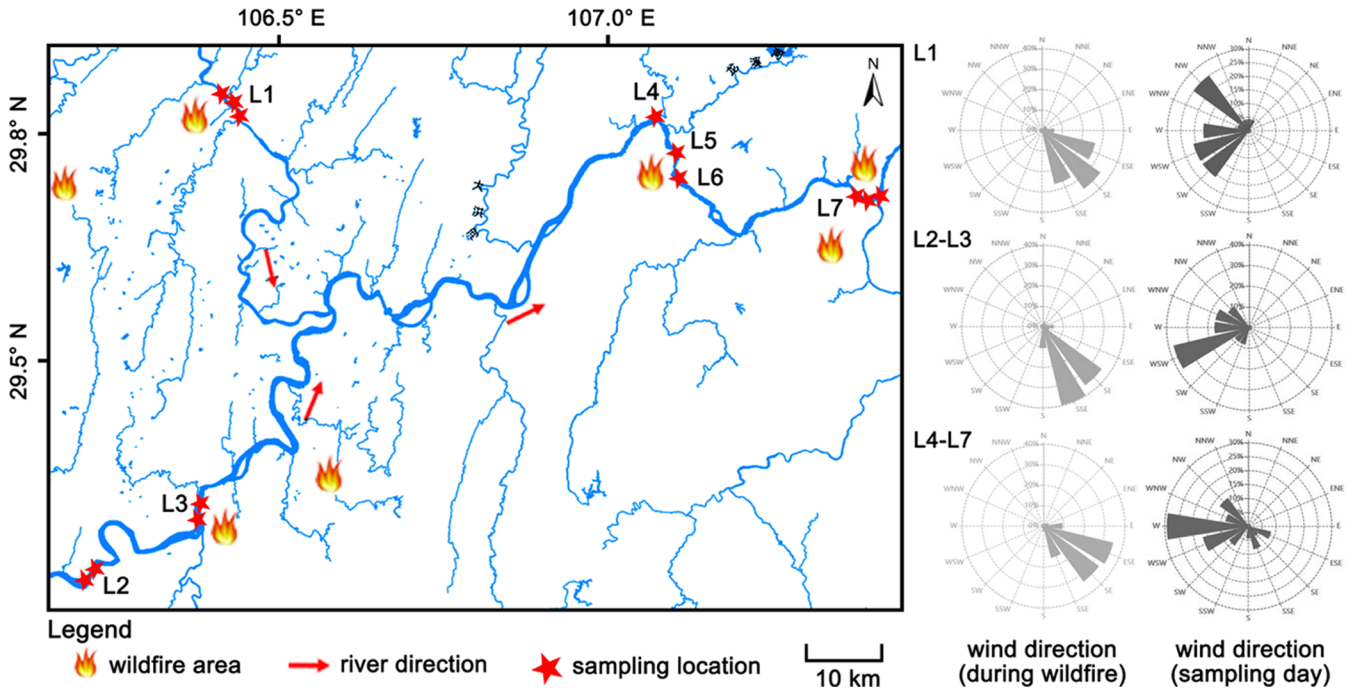


Figure 1. Location of the study area, different wildfire areas, and sampling sites.

determining the postfire export of DOC and DBC.[10,11,15,16] However, in scenarios where there is no rainfall, the immediate impacts of wildfires on aquatic DOM and DBC within a few days postfire remain poorly understood. Apart from soil infiltration and surface runoff,[17] the atmosphere dry deposition of aerosols is recognized as another significant source of DBC in watersheds.[9,18,19] Therefore, it is reasonable to assume that wildfires can have immediate impacts on riverine DOM/DBC via atmospheric deposition of ash.

The aim of this article was to investigate the potential impacts of wind/air deposition of ash on riverine DOM and DBC in the adjacent rivers immediately after wildfires. We collected seven samples a few days after the wildfires from locations near the wildfire-affected areas of the Yangtze and Jialing Rivers located in Chongqing. The riverine DOM and DBC were quantitatively assessed using total organic carbon (TOC) and benzene poly(carboxylic acid) (BPCA) methods, respectively. Additionally, Fourier transform ion cyclotron resonance mass spectrometry (FT-ICR-MS) was employed to characterize the molecular composition of both DOM and DBC. By combining BPCA and FT-ICR-MS analysis, this work aimed to investigate the effects of wildfires on both the quantity and quality of aquatic DOM and DBC.

MATERIALS AND METHODS

Sites and Sampling. Chongqing, situated in southwestern China, covers a total area of 82,400 km^2. The topography in this region is characterized by hills and mountains, and it is intersected by two major rivers, the Yangtze River and the Jialing River, at the center (Figure S1). Since July 2022, Chongqing has been facing an extended period of severe drought marked by minimal precipitation and consistently high temperatures. There have been several forest fires in Chongqing since August 9. On the night of August 17, a forest fire initiated in the mountains and forests within the Fuling District of Chongqing. The fire swiftly propagated to other districts, resulting in a series of fires that endured for 9 days until all wildfires were fully extinguished on August 26.

On August 29, water samples were collected from seven locations along the Yangtze River and the Jialing River (Figures 1 and S2). Table S1 provides information about the sampling locations. Location L2, situated upstream of the main river and far from the nearest wildfire site, should not have been affected by the wildfire event. However, at location L7, burned areas were observed along the riverbank (Figure S2), although no visual differences were observed among the other locations. Notably, there was a notable shift in wind direction from the predominant southeast wind during wildfires to west and west-north winds on the sampling day (Figure 1). The selection of these locations aimed to be as close as possible to the wildfire areas. Replicate samples were taken at approximately 1 km intervals, but some locations had only one or two samples due to geographical constraints. Each water sample (5 L) was directly collected using a 5 L glass bottle from the surface at a depth of 0−20 cm and sent back to the laboratory within 1 day. Water quality parameters including pH, dissolved oxygen (DO), electrical conductivity (EC), oxidation−reduction potential (ORP), and turbidity (TB) were measured in situ using a multiparameter controller on the sampling day. After this, the water samples were subjected to filtration using 0.45 μm membrane filters (Millipore) and then stored at 4 °C for subsequent analyses. In 3 days, DOC was analyzed using a TOC analyzer, and NH_4^+ and NO_3^- were measured using flow injection analysis. The partial filtrate (2 L) was acidified to pH 2 using HCl to inhibit any microorganic activity before DOM extraction using solid-phase extraction.

Extraction of DOM Using Solid-Phase Extraction. DOM was extracted using PPL-resin-based solid-phase extraction (500 mg, 6 mL, Agilent) following the procedure of Dittmar et al.[20] DOM extraction efficiency of PPL cartridges is typically higher than 60%.[21−23] The PPL cartridges were initially rinsed with acidified water and activated with methanol. The filtered water samples were allowed to pass

https://doi.org/10.1021/acs.est.4c00492
Environ. Sci. Technol. XXXX, XXX, XXX−XXX

Table 1. DOC, DBC, DBC/DOC, (B6 + B5)/(B4 + B3) and Basic Quality Parameters at Different Locations

location	DOC (mg/L)	DBC (mg/L)	DBC/DOC	(B6 + B5)/(B4 + B3)	pH	ORP (mV)	NH_4^+ (mg/L)	NO_3^- (mg/L)	EC (μs/cm)	DO (mg/L)	TB (NTU)
L6	5.68	1.09	0.192	7.54	8.05	318	0.077	0.890	331	8.7	7.8
L3	3.36	0.55	0.164	4.90	7.77	89.2	1.824	3.355	347	8.2	8.8
L5	3.32	0.50	0.151	5.66	7.97	348	0.106	0.503	337	8.6	10.6
L1	3.00	0.54	0.181	7.51	7.74	218	0.066	1.329	299	7.4	4.0
L7	1.77	0.22	0.125	0.83	7.77	348	0.105	1.422	354	8.4	7.6
L4	1.66	0.18	0.111	2.80	8.04	333	0.052	0.889	334	8.6	3.9
L2	1.75	0.17	0.095	2.89	7.94	124	0.052	0.890	330	8.2	11.5

through the cartridge at a flow rate of 5 mL min^{-1}. To remove salts, the cartridges were rinsed with 50 mL of acidified water before being dried under N_2. Subsequently, the absorbed DOM was eluted with 10 mL of methanol. The eluted samples were blow-dried with N_2, redissolved in a 1 mL methanol solution, and stored in the dark at −20 °C. For BPCA analysis, each sample from the seven locations was measured. Individual samples from L4, L5, and L6, as well as the second sample from L1, L2, L3, and L7, were employed for FT-ICR-MS analysis.

Quantification of DBC Based on the BPCA Method. DBC concentration was quantified with the BPCA method.[24,25] 0.2 mL of concentrated DOM sample was transferred into a 2 mL Teflon digestion tube. After drying with N_2, the samples were digested with 0.5 mL 65% nitric acid at 170 °C for 9 h. After oxidation, the remaining nitric acid and water were dried under high-purity N_2 at 50 °C. After drying, 1 mL of methanol was added to redissolve the samples, with 10 μg of biphenyl-2′2-dicarboxylic acid as an internal standard.

BPCAs were analyzed using an Agilent 1260 Infinity HPLC system equipped with a photodiode array detector. Samples were eluted at a flow rate of 1 mL/min following gradients from 90% mobile phase A (0.1% phosphoric acid) to 90% mobile phase B (methanol) over 35 min with a Luna phenylhexyl column (5 μm, 4.6 × 250 mm^2). The detailed method is provided in Table S2. The quantification of the BPCAs was carried out with absorbance at 235 nm using the calibration curve of the BPCA standard mixture (Table S3). The standard mixture of the BPCAs was prepared with commercially available BPCAs (1,2,3-B3CA; 1,2,4-B3CA; 1,3,5-B3CA; 1,2,4,5-B4CA; B5CA; B6CA). The calibration curve of 1,2,4,5-B4CA was used to quantify 1,2,3,4-B4CA and 1,2,3,5-B4CA.[26] Moreover, the concentrations of BPCAs were converted to DBC concentrations using an equation proposed by Dittmar.[27]

Ultrahigh-Resolution Mass Spectrometry. A Bruker Solarix 15 T FT-ICR-MS instrument mass spectrometer equipped with a 9.4 T superconducting magnet and an Apollo II electrospray ion source (ESI) was used for mass spectral analysis. DOM samples were injected into an ESI source in negative mode at a rate of 2 μL/min, and each run included 500 scans from 150 to 2000 Da. A total of 128 scans with 2 M word sizes were accumulated to improve the signal-to-noise (S/N) ratio. Mass peaks in the M_w range of 200−800 with S/N greater than 4 and an error value lower than 1.0 ppm were used to assign the correct molecular formula. A molecular formula calculator generated matching formulas according to elemental combinations of $^{12}C_{0-50}$, $^{1}H_{0-150}$, $^{16}O_{0-50}$, $^{14}N_{0-4}$, and $^{32}S_{0-2}$. All elemental formulas must meet the basic chemical criteria: (1) $2 \le H \le (2C + 2)$; (2) $0 \le O \le (C + 2)$; (3) $O/C < 1.2$ and $0.33 \le H/C < 2.3$; (4) $N/C < 0.5$ and $S/C < 0.2$;

The aromaticity index (AI_{mod}) was calculated directly from molecular formulas as (1 + C-0.5O−S-0.5N)/(C-0.5O−S−N). AI_{mod} values ≥ 0.5 and ≥ 0.67 were identified as aromatic and condensed aromatic structures (DBC), respectively.[28] Furthermore, DOM molecules can be categorized into seven groups: (1) condensed aromatics (CAs; $AI_{mod} \ge 0.67$); (2) tannins ($0.67 \le O/C < 1.2$, H/C< 1.5, $AI_{mod} < 0.67$); (3) lignins ($0.1 \le O/C < 0.67$, $H/C < 1.5$, $AI_{mod} < 0.67$); (4) unsaturated hydrocarbons (UHs; $O/C < 0.1$, $H/C < 1.5$, $AI_{mod} < 0.67$); (5) carbohydrates ($0.67 \le O/C < 1.2$, $1.5 \le H/C < 2.3$); (6) proteins ($0.2 \le O/C < 0.67$, $1.5 \le H/C < 2.3$); (7) lipids ($0.67 \le O/C < 1.2$, $1.5 \le H/C < 2.3$); (6) proteins ($O/C < 0.2$, $1.5 \le H/C < 2$).[22,29] The intensity-averaged abundance (wa) parameters such as double-bond equivalence (DBE_{wa}), molecular weight (Mw_{wa}), H/C_{wa}, and O/C_{wa} were calculated based on the assigned formulas of each DOM molecule. Additional details are available in Text S1.

RESULTS AND DISCUSSION

DBC Concentrations and BPCA Composition. Table 1 shows the concentrations of DOC and DBC at sampled sites. The DOC concentrations of samples collected for the seven sites ranged from 1.66 to 5.68 mg/L. The DBC concentrations observed in the studied rivers were 0.17 to 1.09 mg/L, consistent with the global range of riverine DBC concentrations (0.002 to 2.77 mg/L).[30,31] The DBC/DOC ratio in the studied rivers was further calculated and fell within the range of 9.5 to 19.2% (Table 1). Typically, the reported DBC proportions in diverse rivers have exhibited a wide range spanning from 0.1 to 17.5%,[30,31] but the majority remained below 10% (Table 2). In a comprehensive investigation of

Table 2

sampling location	DOC (mg/L)	DBC (mg/L)	DBC/DOC (%)	references
Poudre river	1.56−8.41	0.05−0.34	3.0−6.6	43
Jiulong river	7−18.5	0.06−0.27	0.4−3.2	47
Biomes in North America	0.77−4.01	0.023−0.049	0.4−3.4	12
North America	1.40−2.50	0.04−0.11	2.5−3.8	11
Apalachicola National Estuarine Research Reserve	1.04−7.29	0.007−0.49	0.7−8.0	7
Global rivers	0.67−14.77	0.02−2.06	2.1−15.2	31
Global rivers	−	0.002−2.77	0.1−17.5	30
Chongqing Jialing and Yangtze River	1.66−5.68	0.17−1.09	9.5−19.2	this study

https://doi.org/10.1021/acs.est.4c00492
Environ. Sci. Technol. XXXX, XXX, XXX−XXX

Table 3. Molecular Parameters of the DOM from Different Locations

location	formula	$M_{w,wa}$	C_{wa}	H_{wa}	O_{wa}	N_{wa}	S_{wa}	H/C_{wa}	O/C_{wa}	DBE_{wa}	$(AI_{mod})_{wa}$	$NOSC_{wa}$
L6	5262	439.914	21.15	24.99	9.42	0.210	0.228	1.275	0.464	9.263	0.226	−0.292
L3	5171	412.657	20.12	22.68	8.46	0.286	0.283	1.205	0.445	9.420	0.280	−0.234
L5	5161	416.544	19.82	22.23	9.27	0.259	0.140	1.182	0.484	9.332	0.288	−0.159
L1	5007	452.213	21.42	23.47	10.33	0.261	0.082	1.150	0.498	10.309	0.301	−0.109
L7	4675	456.714	21.57	23.62	10.58	0.189	0.070	1.144	0.500	10.353	0.304	−0.108
L4	4714	454.941	21.51	23.52	10.54	0.180	0.066	1.143	0.499	10.338	0.305	−0.110
L2	4683	457.604	21.58	23.49	10.66	0.176	0.061	1.138	0.503	10.419	0.307	−0.098

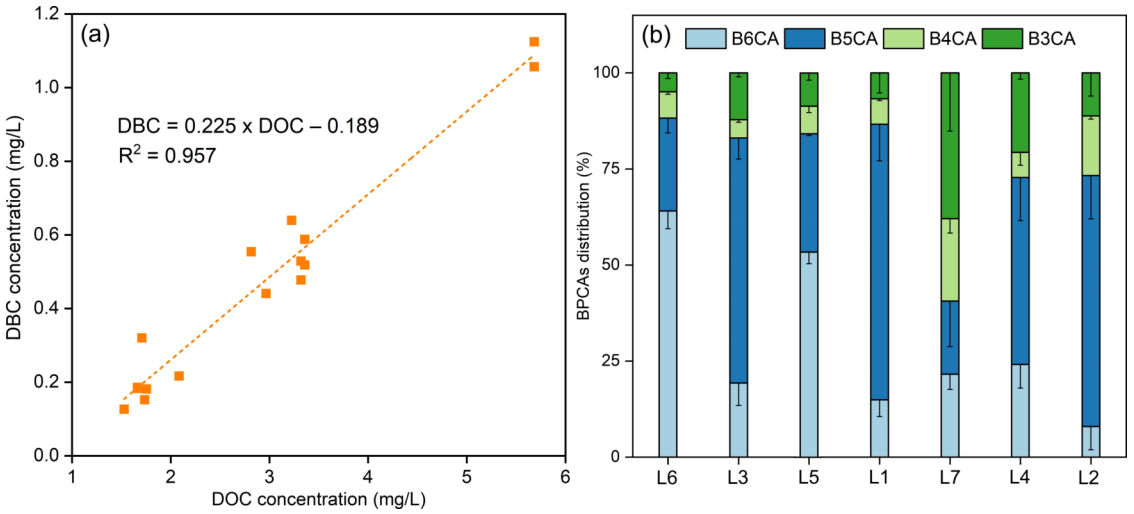


Figure 2. DOC and DBC correlation (a) and BPCA distributions of BPCA-DBC(b) from different locations.

China's major rivers, including the Yangtze, Yellow, Pearl, and Heilongjiang rivers, the DBC concentrations were comparatively lower (0.03−0.34 mg/L).[32] Notably, the Heilongjiang River stood out with the highest DOC concentrations (5.62−12.86 mg/L) but much lower DBC concentrations (0.07−0.09 mg/L) compared to other rivers.[32] It reflected a relatively lower input of DBC in the Heilongjiang region, attributed to its less populated and colder northern climate with fewer landscape fires, in contrast to the more densely populated and warmer southern regions that experience more biomass burning.[32] At location L2 that was obtained from upstream of the main river (Yangtze River) and far from the nearest wildfire site (21 km), the lowest DBC concentration (0.17 mg/L) was detected, indicating that it might be nonimpacted. At locations L1, L3, L5, and L6, where higher levels of DOC were recorded, correspondingly, elevated concentrations of DBC were also observed. These specific sites were also situated in close proximity to points of wildfire occurrence (2.5−5.2 km), suggesting a possible impact from the wildfire events. Typically, the hydrology during rain and storm events was considered the primary factor driving the input of DBC from wildfire-affected lands.[14,15,33] However, due to the absence of rainfall between the wildfire event and the sampling day (Figure S3), the hydrology was not the reason for riverine DBC input. Meanwhile, the daily average concentrations of atmospheric OC and BC experienced multiple peaks on wildfire days, followed by a decline to their lowest levels on August 27th, after the wildfires (Figure S4).[25,34] Thus, the atmospheric ash deposition might be a significant contributor to riverine DBC. Notably, all samples were collected 3 days postfire, which may have resulted in changes in concentrations or compositions during the initial days via potential biogeochemical processes.[35−37] Especially, despite its proximity to two wildfire areas (1.1−2.3 km), L7 exhibited relatively low levels of DOC and DBC. Unlike other locations (e.g., L1, L5, and L6) located downwind on the sampling day, L7 was parallel to the prevailing wind direction (Figure 1). Consequently, the wind patterns on the sampling day might have facilitated ash deposition in adjacent rivers. Therefore, it is reasonable to infer that locations L2, L4, and L7 were nonimpacted or less impacted by wildfires, while L1, L3, L5, and L6 were highly impacted, potentially due to atmospheric deposition of ash facilitated by wind.

Although noticeable variability was observed across all sites, a strong correlation existed between DBC and DOC concentrations, as evidenced by a high correlation coefficient ($R^2 = 0.957$, $p < 0.05$) (Figure 2a). A high correlation between DOC and DBC in global rivers has also been documented, indicating that the mobilization of riverine DBC and DOC was mechanistically coupled.[30] A diminished correlation between DBC and DOC concentrations has been observed in some rivers, including the Amazon River,[38] and China's major rivers.[32] This phenomenon was attributed to the distinct impacts of environmental factors, such as soil properties, temperature, rainfall patterns, and aerosol deposition, on the mobilization of DOC and DBC.[38] Within the context of the studied rivers, the occurrence of sudden wildfire incidents likely intensified the correlation between DBC and DOC. Meanwhile, the strong coupling between atmospheric OC and BC during wildfires in this area (Figure S4) also indicated that riverine DOC and DBC were likely sourced from ash deposition. Studies have also found a coupling of DOC and certain water quality parameters, such as turbidity,[13,14] suggesting that DOM likely originated from mobilized surface

https://doi.org/10.1021/acs.est.4c00492
Environ. Sci. Technol. XXXX, XXX, XXX−XXX

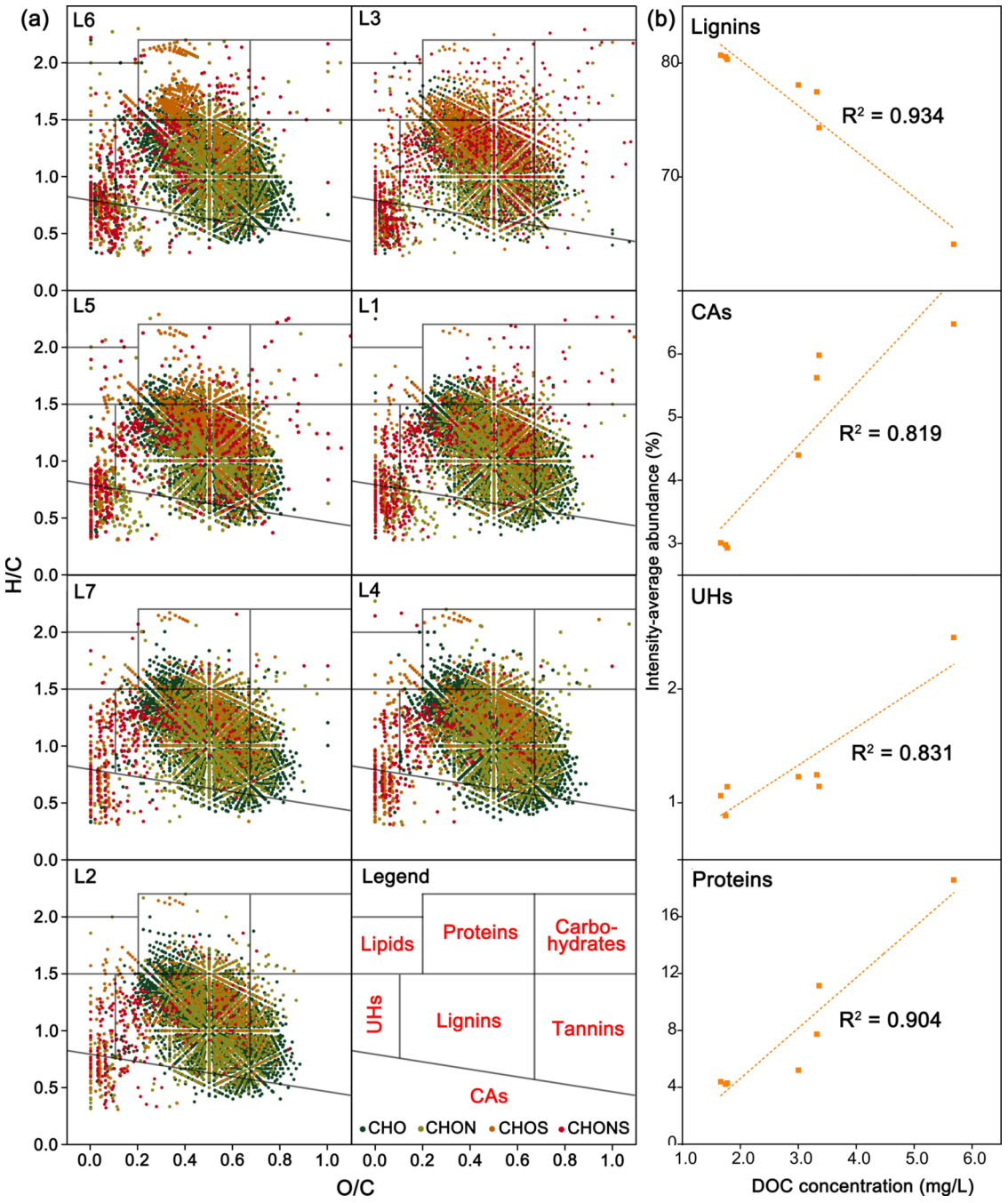


Figure 3. (a) van Krevelen diagrams of DOM from different locations. (b) Correlations between DOC concentration and intensity-averaged abundance of lignins, CAs, UHs, and proteins.

materials due to postfire runoff and erosion.[13] Salinity has been widely documented to negatively correlate with aquatic DBC via affecting the physical and photochemical processes of DBC.[39,40] Notably, there were no significant correlations between tested water quality parameters and DOC or DBC (Figures S5 and S6), indicating that their dynamics were decoupled from DOC or DBC. The decoupling indicated that these parameters might be less influenced by ash deposition, and riverine DOC and DBC deposited immediately after wildfires were less affected by water chemistry.

The BPCA distributions of DBC in Chongqing Jialing and the Yangtze River are shown in Figure 2b. Through nitric oxidation, the formation of highly substituted BPCAs (i.e., B6CA and B5CA) occurs when aromatic rings are surrounded by multiple aromatic rings, while poorly substituted BPCAs (i.e., B3CA) form when aromatic rings are surrounded by only a few aromatic rings.[26,41] In comparison to B3CA and B4CA, B5CA and B6CA serve as indicators of DBC with a more pronounced aromatic structure and increased condensation.[28,42] In the studied rivers, the predominant BPCAs were B5CAs (19.0, 71.7%) and B6CA (8.0, 64.1%), followed by B3CA (4.9, 37.9%) and B4CA (4.8−21.5%). The degree of condensed aromaticity of DBC was further estimated using the ratio of (B6CA + B5CA)/(B4CA + B3CA) (Table 1). The BPCA ratio is associated with pyrogenic source material and prevailing environmental conditions.[43] For example, a higher

https://doi.org/10.1021/acs.est.4c00492
Environ. Sci. Technol. XXXX, XXX, XXX−XXX

BPCA ratio, indicating more condensed aromatic DBC, is linked to DBC originating from watersheds affected by wildfires,[43] whereas DBC originated from urban dust and fossil fuel combustion exhibits a lower ratio.[11,25] Furthermore, environmental factors such as prolonged solar radiation exposure[28] and elevated salinity levels[44,45] are linked to relatively low BPCA ratios. In the studied rivers, the BPCA ratios ranged from 0.83 to 7.54, surpassing the reported values in rivers of the Bohai Rim region (0.40−0.55),[45,46] and surface waters of Alpine, Antarctic, and Arctic regions (0.02−0.77).[25] Thus, it is reasonable to infer that the abrupt wildfires played a role in introducing riverine DBC with high condensed aromaticity. Given that the samples were collected shortly after the wildfires, the pronounced condensed aromaticity of DBC appeared to be minimally influenced by the water conditions or potential photodegradation from solar radiation. Notably, at locations L1, L3, L5, and L6 with high DBC concentrations and DBC/DOC, there was a corresponding high degree of polycondensation of DBC (Table 1).

General Molecular Characterization of DOM Using FT-ICR-MS. Using FT-ICR-MS, a total of 10,280 unique molecular formulas were identified across all DOM samples, of which 19.7% were common/shared, while 41.7% were unique and distinct from each other. Generally, each DOM sample contained a range of 4675−5262 unique molecular formulas. The CHO compounds were the most abundant species of DOM, constituting 2053−2575 (39.7−54.8%) of the total assigned molecular formulas, and 56.0−82.0% of the intensity-averaged abundance. The CHO compounds contain typical products resulting from incomplete combustion,[48] have been widely detected in the ESI negative mode and identified in aquatic DOM.[49,50] Notably, DOM samples from L1, L3, L5, and L6 with higher DOC concentrations contained more molecular formulas (5007−5262 vs 4675−4714). However, the CHO compounds from locations L1, L3, L5, and L6 contained lower number of formulas (2052−2381, 39.7−47.0%) compared to locations L2, L4, and L7 (2528−2575, 54.1−54.8%). As a result, wildfires potentially enhanced the molecular complexity and diversity of DOM but reduced the diversity and abundance of CHO compounds in the studied rivers.[49,51] Biomass combustion that introduces oxygen to a molecule can reduce its volatility and hydrophobicity, subsequently increasing its water solubility.[52] The high degree of oxidation was indicative of the presence of acidic functional groups, such as hydroxyl and carboxylic acids.[22] The majority of detected CHO compounds were heavily oxidized (O ≥ 10), constituting 56.3−66.9% of the total number of CHO formulas. However, locations L1, L3, L5, and L6 contained more low-oxygen-containing CHO species (O < 10) (Figure 4), indicating that the incomplete combustion of biomass caused by limited oxygen availability resulted in the formation of low-oxygen-containing CHO compounds.

As shown in the van Krevelen diagram (Figure 3a), the dominant composition among DOM molecules was lignins, followed by tannins, protein-like compounds, CAs, UHs, carbohydrates, and lipids. Notably, in locations highly impacted by wildfires (L1, L3, L5, and L6), there were lower abundances of lignins (58.3−64.5 vs 66.9−67.5%) and tannins (8.9−12.4 vs 12.6−13.4%), yet higher abundances of protein-like compounds (8.8−13.4 vs 7.9−8.2%), CAs (11.2−11.6 vs 8.6−8.8%), and UHs (2.2−5.1 vs 1.8−2.2%) compared to L2, L4, and L7. Additionally, the intensity-averaged abundances of protein-like compounds, CAs, and UHs exhibited a strong positive correlation with the DOC concentration of DOM (R^2 = 0.904, 0.831, and 819, respectively), whereas the abundance of lignins was negatively correlated with the DOC concentration (R^2 = 0.934) (Figure 3b). It revealed a strong association between wildfires and the increase in protein-like compounds, CAs, and UHs, along with a decrease in lignins. Chen et al. also reported that the identified formulas in DOM from burned materials contained a higher proportion of proteins (7.26−8.14 vs 0.59−5.23%) and CAs (20.4−27.1 vs 3.04−4.67%) compared to unburned material.[49] During wildfires, the conversion of plant-derived lignin/phenol to (poly)aromatic carbon accompanied by the loss of side chains was extensively documented, and as a result, there was an increase in aromaticity and a decrease in molecular weight.[49,51] However, in this study, the aromaticity ($(AI_{mod})_{wa}$) and double-bond equivalence (DBE_{wa}) were negatively correlated with the DOC (Table S5). It was attributed to the production of low-aromaticity protein-like compounds, as evidenced by negative correlations between $(AI_{mod})_{wa}$ and DBE_{wa} of protein-like compounds and DOC (Table S5). Notably, there was a strong negative correlation between DOC and the redox potential of DOM molecules ($NOSC_{wa}$) (Table S5), indicating that DOM with a high DOC was more reductive. The decoupling of DOC and ORP of water (Figure S5) indicated that other factors may affect the redox state of the aquatic environment, independent of the DOM potentially derived from ash deposition.

Nitrogen-Containing DOM Release under Wildfires Impacts. The CHON compounds were the second most frequently detected species, constituting 1484−1776 formulas, accounting for 31.7−35.5% of the total formulas (Table S6). A higher number of CHON formulas were detected in locations L1, L3, L5, and L6 relative to locations L2, L4, and L7 (1678−1776 vs 1484−1523). Chen et al. also reported that DOM from burned materials contained a higher percentage of CHON compounds compared to unburned materials (33.5−45.3% versus 5.56−14.1%).[49] It was related to the creation of heterocyclic nitrogen compounds during the wildfires, leading to the preferential fixation of nitrogen over carbon.[53] As reported, wildfires, particularly under flaming conditions, are known to release CHON species like nitrogen oxides (NOx) and nitrous acid (HONO).[54,55] These short-lived compounds may subsequently undergo transformations into longer-lived compounds, such as peroxyacetyl nitrate (PAN), nitric acid (HNO_3) organic nitrates, or other organic nitrogen-containing compounds.[54] The CHON compounds were classified into 9 subgroups according to their N and O numbers, including N_1O_x, N_2O_x, N_3O_x, and N_4O_x. According to previous studies, an amine or nitro group typically has an O/N ratio lower than 3, while the high O/N ratios (≥3) of these formulas allow an assignment of one nitro ($-NO_2$) or nitrooxy ($-ONO_2$) group, and PAN exhibits an O/N ratio of 5 or higher.[56,57] In this study, it was found that 81.0−94.0% of the CHON compounds had an O/N ratio of ≥ 3, and 71.4−86.1% had an O/N ratio of ≥ 5. Hence, it can be inferred that most of CHON compounds were likely PAN compounds, exhibiting quite high polarity and water solubility.[57] Notably, fewer formulas with O/N ≥ 3 were found for L1, L3, L5, and L6 compared to L2, L4, and L7 (81.0−89.4 vs 92.6−94.0%). The difference was attributed to the N_1O_x species, implying that wildfires contributed to the formation of more reduced N_1O_x species. However, the abundances of both O/N < 3 and O/N ≥ 3 of N_2O_x species were higher at locations highly affected by wildfires. Moreover,

https://doi.org/10.1021/acs.est.4c00492
Environ. Sci. Technol. XXXX, XXX, XXX−XXX

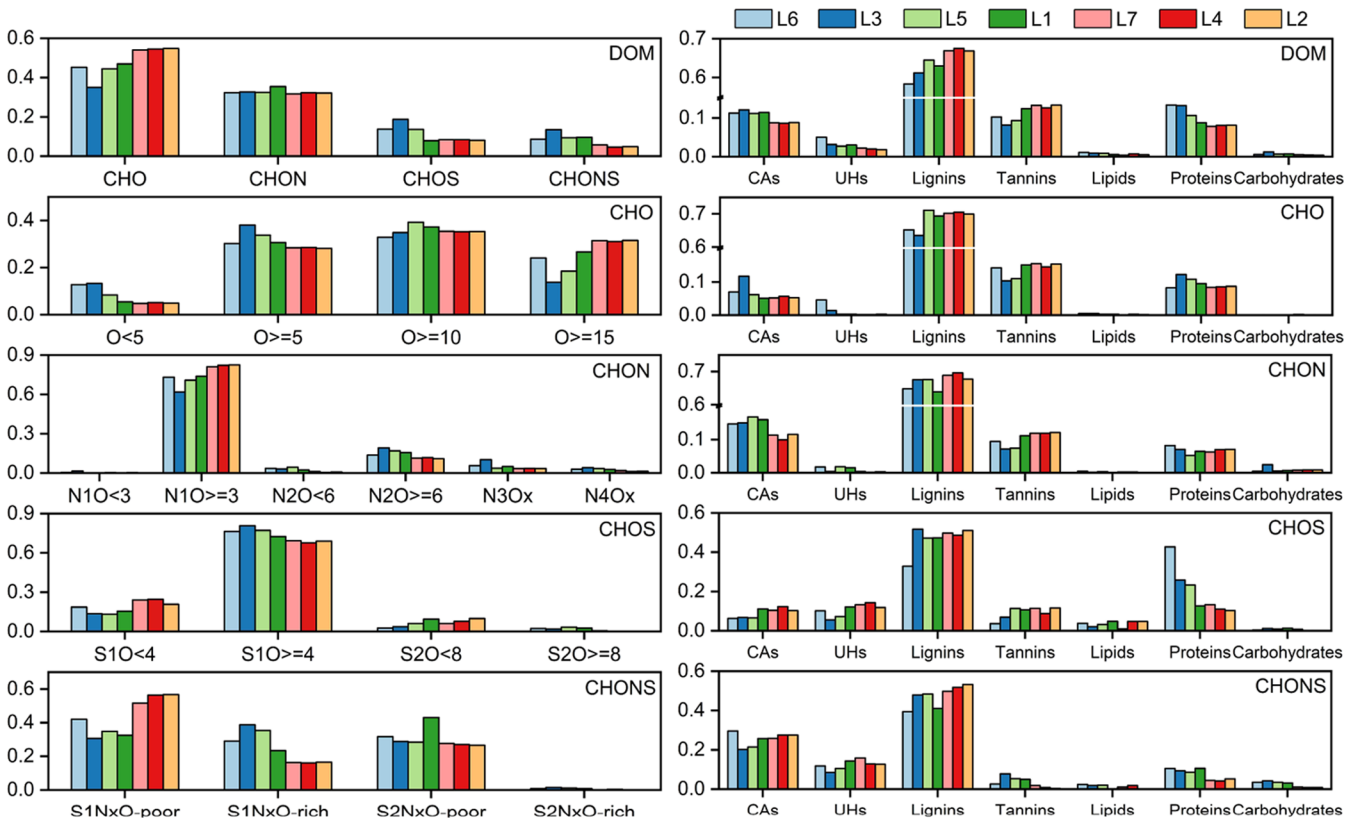


Figure 4. (Left) Formula number proportion of elemental composition-based subgroups for respective DOM, CHO, CHON, CHOS, and CHONS compounds; (Right) Formula number proportion of seven compound classifications for respective DOM, CHO, CHON, CHOS, and CHONS compounds.

the CHON compounds at these locations contained more N_3O_x and N_4O_x species, almost all of which had O/N ratios of less than 3. Therefore, wildfires were strongly correlated with the generation of reduced nitrogen compounds, particularly those containing a higher number of nitrogen atoms. As reported, these reduced nitrogen compounds might be mainly associated with alkyl amino and alkyl nitrile as well as heterocyclic aromatic rings with a single N atom.[58,59] Notably, the relative abundances of condensed aromatics within the CHON compounds were higher in locations L1, L3, L5, and L6 in contrast to L2, L4, and L7 (13.2−16.9 vs 10.0−11.6%). The elevated proportion of CAs in these locations was linked to the reduced abundance of oxygen-rich tannins and/or lignins, while no notable alterations were observed in the proportions of UHs, proteins, and carbohydrates.

CAs can be divided into CHO, CHON, CHOS, and CHONS species based on the presence of heteroatoms (Figures S10 and S11). Only 20.9−35.9% of the formulas of CAs were characterized as CHO species, while the other 64.1−79.1% contained CHO and one or more heteroatoms such as N and S. The CHON and CHOS species of CAs are also known as dissolved black nitrogen (DBN) and dissolved black sulfur (DBS), originating from the incorporating of organic nitrogen and sulfur into DBC during biomass burning.[12,60] Similar to CHO species, the DBS formula numbers exhibited no significant discrepancy between locations impacted and unimpacted by wildfires (Figure S10). Remarkably, DBN constituted the highest abundance, accounting for 37.3−49.7% (152−286 formulas) of the total CAs. More DBN formulas (227−286) were detected in locations highly affected by wildfires compared to other locations (152−175). A study also revealed that the levels of DBN within DOM from burned biomass were 5.8 to 21.2 times higher than those found in unburned biomass,[49] highlighting the significant role of wildfires in the formation of DBN. Additionally, relatively high formulas (104−142 vs 60−70) of the CHONS species were also detected in locations highly affected by wildfires. The CHONS species of CAs also partially explained the increased CHONS compounds within the DOM under the impacts of wildfires (Figure 4).

Sulfur-Containing DOM Release under Wildfires Impacts. The CHOS (379−979 formulas) and CHONS compounds (218−703 formulas) accounted for 7.9−18.8% and 4.6−13.5% of the total identified compounds in studied rivers, respectively (Figure 4). Notably, locations highly affected by wildfires (L3, L5, and L6) exhibited a greater number of CHOS formulas, ranging from 704 to 979, compared to other locations with 379−399 formulas. The intensity-averaged abundance of CHOS was also diverse among the tested samples, with higher percentages of CHOS compounds in locations L1, L3, L5, and L6 compared to other locations (4.1−4.6 vs 8.9−18.1%). Similarly, locations L1, L3, L5, and L6 contained more CHONS species relative to locations L2, L4, and L7 (456−703 vs 218−271, 8.7−13.5 vs 4.6−5.8%). Sulfur that can be sourced from both the soil and the atmosphere in the form of SO_2 and sulfate is a vital nutrient for plants as a part of biomass such as disulfide bonds in plant proteins and a range of other sulfur-containing (as well as oxygen- and nitrogen-containing) compounds.[61] Although there have been limited studies on the specific contribution of wildfires to the release of sulfur-containing compounds into rivers, existing research suggested that biomass burning plays

https://doi.org/10.1021/acs.est.4c00492
Environ. Sci. Technol. XXXX, XXX, XXX−XXX

an important role in the emission of sulfur-containing compounds to the atmosphere.[48,57,61] In this study, due to atmospheric deposition, wildfires might contribute to the input of abundant sulfur-containing compounds into the rivers.

The identified CHOS formulas were divided into four subgroups based on their S and O numbers (Figure 4). The O-rich CHOS fraction (O/S ≥ 4) typically contains organic sulfates ($-OSO_3H$) or sulfonates ($-SO_3$), whereas when O/S < 4, it suggests the presence of less oxidized sulfur functional groups such as sulfones ($-SO_2-$), sulfonic acid ($-SO_3H$), sulfinic acid ($-SO_2H$), or sulfides.[48] 67.7−80.6% of CHOS formulas was identified as O-rich species, with intensity-averaged abundance of 57.3−91.2%, indicating that most of the CHOS compounds detected by negative ESI sources likely contained sulfate ($-OSO_3H$) or sulfonate ($-SO_3$) groups. Based on N, S, and O atoms, CHONS compounds were classified into O-poor ((4S + 3N)/O < 1) and O-rich ((4S + 3N)/O ≥ 1) compounds. The O-rich CHONS compounds are likely nitrooxy−organosulfates containing nitrate ($-ONO_2$) groups, while O-poor species might contain reduced sulfur (e.g., sulfones or aromatic sulfur) or reduced nitrogen functional groups (e.g., amine or heterocyclic aromatics) due to their insufficient oxygen content in the molecular formulas.[57] As the majority of CHONS compounds were O-poor compounds that accounted for 59.6−83.5% of total formulas, these compounds mostly contained reduced nitrogen and sulfur functional groups. Ditto et al. also indicated that the CHONS compounds emitted from a forest fire contained abundant aromatic nitrogen, amine, and imine, as well as sulfide, aromatic sulfur, and sulfone functional groups.[61]

Although there were more both O-poor and O-rich sulfur-containing formulas observed in locations affected by wildfires, it is important to note that wildfires potentially increased the relative abundance of both the O-rich CHOS and CHONS compounds, as indicated by the more abundant of the O-rich species, especially those containing one S atom, in locations L1, L3, L5, and L6 compared to L2, L4, and L6 (Figure 4). In addition, for both CHOS and CHONS compounds, the relatively high O/C_{wa} and H/C_{wa} but low DBE/C_{wa} and $(AI_{mod})_{wa}$ were found in locations highly affected by wildfires (Table S6), indicating the formation of oxidative, aliphatic, saturated, and low-aromatic S-containing compounds under the impacts of wildfires. In locations L1, L3, L5, and L6, there was an increased presence of sulfur-containing lignins and protein-like compound formulas compared to other locations. In terms of relative abundance, higher percentages of protein-like substances were detected in these wildfire-impacted areas (Figure 3). In particular, at location L6, the abundance of protein-like CHOS compound formulas and their intensity-averaged abundance were notably high, reaching 44.1% and 80.7%, respectively.

■ ENVIRONMENTAL IMPLICATIONS

The globally increasing extreme high temperatures in recent decades have led to a rise in wildfires, posing a significant threat to global ecosystems and impacting ecological security. Wildfires typically transport pyrogenic organic matter and nutrients to rivers through interactions with the hydrological cycle.[13] Our study highlighted that atmosphere ash deposition was likely responsible for the high levels of DOC and DBC in locations close to the wildfire areas, but wind direction on the sampling day also played a role. The coupling of DOC and DBC, along with the decoupling of DOC and other water quality parameters, indicated that ash deposition might not significantly affect water chemistry but might contribute to aquatic DOM/DBC dynamics. Besides, DOM/DBC sourced from abrupt wildfires was not significantly affected by the original water chemistry in the short term. Based on the molecular composition analysis, the majority of DBC was DBN which mainly accounted for the increase in N-containing DOM. Moreover, sulfur-containing substances in plants were integrated into the pyrogenic OM during wildfires, contributing to the generation of abundant recalcitrant DOS.

Such pyrogenic DBN and DOS may be more recalcitrant to photodegradation and biotransformation than the bulk DOM and thus accumulate in aquatic systems.[62,63] These substances may further influence various aspects of water chemistry, such as pH[64,65] and redox potential,[66] as well as the biogeochemical cycling of trace metals.[67] For instance, a strong association was observed between DBN/DOS abundance and the reductive activity of DOM molecules. It suggested that these pyrogenic DOM have the potential to alter the water redox potential and element cycling, despite no relationships being observed in the short term. In addition, the high photoactivity of pyrogenic DOM in the DOM pool may contribute to contaminant photodegradation via photosensitivity processes.[68,69] The appearance of riverine pyrogenic DOM may modify microbial loop dynamics, shaping the structure and function of riverine microbial communities, including processes related to nitrogen and sulfur cycling.[70−72] Therefore, the wildfire-derived pyrogenic DOM may potentially impact water quality and environmental biogeochemistry in various ways. Further research is needed to understand their specific effects and contributions to ecosystem processes in the long term. A comprehensive understanding of the fate and transformation, including aspects such as photoreactivity and bioavailability, of wildfires-derived pyrogenic DOM is crucial for a more thorough comprehension of its environmental implications.

■ ASSOCIATED CONTENT

Supporting Information

The Supporting Information is available free of charge at https://pubs.acs.org/doi/10.1021/acs.est.4c00492.

Analysis of FT-ICR-MS data (Text S1); Chongqing's map and location on the map of China (Figure S1); pictures of several sampling locations (Figure S2); amount of precipitation (mm) in Chongqing before and after wildfire events (Figure S3); daily average atmospheric OC and BC concentrations in Chongqing before and after wildfire events (Figure S4); correlation analysis of tested water quality parameters and DOC (Figure S5); correlation analysis of tested water quality parameters and DBC (Figure S6); full spectra of (−) ESI FT-ICR-MS of PPL-DOMs at different locations (Figure S7); intensity-averaged proportion of elemental composition-based subgroups for respective DOM, CHO, CHON, CHOS, and CHONS compounds, and intensity-averaged proportion of seven compound classifications for respective DOM, CHO, CHON, CHOS, and CHONS compounds (Figure S8); correction between DOC concentration of DOM and intensity-averaged abundance of tannins, carbohydrates, and lipids (Figure S9); formula number of elemental composition-based subgroups for respective DOM, CAs, UHs, lignins, tannins, lipids, proteins, and carbohydrates

https://doi.org/10.1021/acs.est.4c00492
Environ. Sci. Technol. XXXX, XXX, XXX−XXX

(Figure S10); formula number proportion of elemental composition-based subgroups for respective DOM, CAs, UHs, lignins, tannins, lipids, proteins, and carbohydrates (Figure S11); intensity-averaged proportion of elemental composition-based subgroups for respective DOM, CAs, UHs, lignins, tannins, lipids, proteins, and carbohydrates (Figure S12); information about sampling locations (Table S1); mobile phase mixing gradients (Table S2); calibration curves of BPCAs target compounds (Table S3); molecular parameters of DOM with different molecular compositions (Table S4); correlations between DOC and molecular parameters of DOM with different molecular compositions (Table S5); molecular parameters of DOM with different elemental compositions (Table S6); and correlations between DOC and molecular parameters of DOM with different elemental compositions (Table S7) (PDF)

AUTHOR INFORMATION

Corresponding Authors

Zhongbo Zhou − *College of Resources and Environment, Southwest University, Chongqing 400715, China*; Email: zhouzb86@swu.edu.cn

Gang Liu − *Key Laboratory of Drinking Water Science and Technology, Research Centre for Eco-Environmental Sciences, Chinese Academy of Sciences, Beijing 100085, China; Section of Sanitary Engineering, Department of Water Management, Faculty of Civil Engineering and Geosciences, Delft University of Technology, 2628 CN Delft, The Netherlands; University of Chinese Academy of Sciences, Beijing 101408, China*; orcid.org/0000-0002-4008-9017; Email: g.liu-1@tudelft.nl, gliu@rcees.ac.cn

Authors

Yanghui Xu − *Key Laboratory of Drinking Water Science and Technology, Research Centre for Eco-Environmental Sciences, Chinese Academy of Sciences, Beijing 100085, China; Section of Sanitary Engineering, Department of Water Management, Faculty of Civil Engineering and Geosciences, Delft University of Technology, 2628 CN Delft, The Netherlands*

Xintu Wang − *Key Laboratory of Drinking Water Science and Technology, Research Centre for Eco-Environmental Sciences, Chinese Academy of Sciences, Beijing 100085, China; College of Environmental Science and Engineering, Guilin University of Technology, Guangxi 541004, China*

Qin Ou − *Key Laboratory of Drinking Water Science and Technology, Research Centre for Eco-Environmental Sciences, Chinese Academy of Sciences, Beijing 100085, China; Section of Sanitary Engineering, Department of Water Management, Faculty of Civil Engineering and Geosciences, Delft University of Technology, 2628 CN Delft, The Netherlands*

Jan Peter van der Hoek − *Section of Sanitary Engineering, Department of Water Management, Faculty of Civil Engineering and Geosciences, Delft University of Technology, 2628 CN Delft, The Netherlands; Waternet, Department Research & Innovation, 1090 GJ Amsterdam, The Netherlands*; orcid.org/0000-0002-0674-388X

Complete contact information is available at: https://pubs.acs.org/10.1021/acs.est.4c00492

Author Contributions

[▽]Y.X. and X.W. contributed equally to this work.

Notes

The authors declare no competing financial interest.

ACKNOWLEDGMENTS

The authors acknowledge the support from National Natural Science Foundation of China (52388101; 32161143031). Dr. Zhongbo Zhou also thanks the support from the merit of Chongqing Bayu Young Scholar (YS2021015) in China.

REFERENCES

(1) Flannigan, M. D.; Krawchuk, M. A.; de Groot, W. J.; Wotton, B. M.; Gowman, L. M. Implications of changing climate for global wildland fire. *Int. J. Wildland Fire* **2009**, *18* (5), 483−507, DOI: 10.1071/WF08187.

(2) Krawchuk, M. A.; Moritz, M. A.; Parisien, M.-A.; Van Dorn, J.; Hayhoe, K. Global pyrogeography: the current and future distribution of wildfire. *PLoS One* **2009**, *4* (4), No. e5102.

(3) Masiello, C. A. New directions in black carbon organic geochemistry. *Mar. Chem.* **2004**, *92* (1−4), 201−213.

(4) Lian, F.; Xing, B. Black carbon (biochar) in water/soil environments: molecular structure, sorption, stability, and potential risk. *Environ. Sci. Technol.* **2017**, *51* (23), 13517−13532.

(5) Moody, J. A.; Martin, D. A. Initial hydrologic and geomorphic response following a wildfire in the Colorado Front Range. *Earth Surf. Processes Landforms* **2001**, *26* (10), 1049−1070, DOI: 10.1002/esp.253.

(6) Lane, P. N.; Sheridan, G. J.; Noske, P. J. Changes in sediment loads and discharge from small mountain catchments following wildfire in south eastern Australia. *J. Hydrol.* **2006**, *331* (3−4), 495−510.

(7) Dittmar, T.; Paeng, J.; Gihring, T. M.; Suryaputra, I. G. N. A.; Huettel, M. Discharge of dissolved black carbon from a fire-affected intertidal system. *Limnol. Oceanogr.* **2012**, *57* (4), 1171−1181.

(8) Wagner, S.; Jaffé, R.; Stubbins, A. Dissolved black carbon in aquatic ecosystems. *Limnol. Oceanogr. Lett.* **2018**, *3* (3), 168−185.

(9) Bao, H.; Niggemann, J.; Luo, L.; Dittmar, T.; Kao, S. J. Aerosols as a source of dissolved black carbon to the ocean. *Nat. Commun.* **2017**, *8* (1), No. 510.

(10) Dittmar, T.; de Rezende, C. E.; Manecki, M.; Niggemann, J.; Ovalle, A. R. C.; Stubbins, A.; Bernardes, M. C. Continuous flux of dissolved black carbon from a vanished tropical forest biome. *Nat. Geosci.* **2012**, *5* (9), 618−622, DOI: 10.1038/ngeo1541.

(11) Ding, Y.; Yamashita, Y.; Dodds, W. K.; Jaffe, R. Dissolved black carbon in grassland streams: is there an effect of recent fire history? *Chemosphere* **2013**, *90* (10), 2557−2562.

(12) Ding, Y.; Watanabe, A.; Jaffé, R. Dissolved black nitrogen (DBN) in freshwater environments. *Org. Geochem.* **2014**, *68*, 1−4.

(13) Roebuck, J. A.; Bladon, K. D.; Donahue, D.; Graham, E. B.; Grieger, S.; Morgenstern, K.; Norwood, M. J.; Wampler, K. A.; Erkert, L.; Renteria, L.; Danczak, R.; Fricke, S.; Myers-Pigg, A. N. Spatiotemporal Controls on the Delivery of Dissolved Organic Matter to Streams Following a Wildfire. *Geophys. Res. Lett.* **2022**, *49* (16), No. e2022GL099535.

(14) Uzun, H.; Dahlgren, R. A.; Olivares, C.; Erdem, C. U.; Karanfil, T.; Chow, A. T. Two years of post-wildfire impacts on dissolved organic matter, nitrogen, and precursors of disinfection by-products in California stream waters. *Water Res.* **2020**, *181*, No. 115891.

(15) Barton, R.; Richardson, C. M.; Pae, E.; Montalvo, M. S.; Redmond, M.; Zimmer, M. A.; Wagner, S. Hydrology, rather than wildfire burn extent, determines post-fire organic and black carbon export from mountain rivers in central coastal California. *Limnol. Oceanogr. Lett.* **2024**, *9* (1), 70−80.

(16) Myers-Pigg, A. N.; Louchouarn, P.; Amon, R. M. W.; Prokushkin, A.; Pierce, K.; Rubtsov, A. Labile pyrogenic dissolved organic carbon in major Siberian Arctic rivers: Implications for wildfire-stream metabolic linkages. *Geophys. Res. Lett.* **2015**, *42* (2), 377−385.

(17) Chen, Y.; Sun, K.; Wang, Z.; Zhang, E.; Yang, Y.; Xing, B. Analytical methods, molecular structures and biogeochemical behaviors of dissolved black carbon. *Carbon Res.* **2022**, *1* (1), No. 23, DOI: 10.1007/s44246-022-00022-4.
(18) Jones, M. W.; Quine, T. A.; de Rezende, C. E.; Dittmar, T.; Johnson, B.; Manecki, M.; Marques, J. S. J.; de Aragão, L. E. O. C. Do Regional Aerosols Contribute to the Riverine Export of Dissolved Black Carbon? *J. Geophys. Res.: Biogeosci.* **2017**, *122* (11), 2925−2938.
(19) Wozniak, A. S.; Bauer, J. E.; Dickhut, R. M.Fossil and contemporary aerosol particulate organic carbon in the eastern United States: Implications for deposition and inputs to watersheds *Global Biogeochem. Cycles* 2011; Vol. *25* 2 DOI: 10.1029/2010GB003855.
(20) Dittmar, T.; Koch, B.; Hertkorn, N.; Kattner, G. A simple and efficient method for the solid-phase extraction of dissolved organic matter (SPE-DOM) from seawater. *Limnol. Oceanogr.: Methods* **2008**, *6* (6), 230−235.
(21) Xu, W.; Gao, Q.; He, C.; Shi, Q.; Hou, Z. Q.; Zhao, H. Z. Using ESI FT-ICR MS to Characterize Dissolved Organic Matter in Salt Lakes with Different Salinity. *Environ. Sci. Technol.* **2020**, *54* (20), 12929−12937.
(22) Zhao, P.; Du, Z.; Fu, Q.; Ai, J.; Hu, A.; Wang, D.; Zhang, W. Molecular composition and chemodiversity of dissolved organic matter in wastewater sludge via Fourier transform ion cyclotron resonance mass spectrometry: Effects of extraction methods and electrospray ionization modes. *Water Res.* **2023**, *232*, No. 119687.
(23) Green, N. W.; Perdue, E. M.; Aiken, G. R.; Butler, K. D.; Chen, H.; Dittmar, T.; Niggemann, J.; Stubbins, A. An intercomparison of three methods for the large-scale isolation of oceanic dissolved organic matter. *Mar. Chem.* **2014**, *161*, 14−19.
(24) Geng, X.; Zhong, G.; Liu, J.; Sun, Y.; Yi, X.; Bong, C. W.; Zakaria, M. P.; Gustafsson, Ö.; Ouyang, Y.; Zhang, G. Year-Round Measurements of Dissolved Black Carbon in Coastal Southeast Asia Aerosols: Rethinking Its Atmospheric Deposition in the Ocean. *J. Geophys. Res.: Atmos.* **2021**, *126* (18), No. e2021JD034590.
(25) Khan, A. L.; Wagner, S.; Jaffe, R.; Xian, P.; Williams, M.; Armstrong, R.; McKnight, D. Dissolved black carbon in the global cryosphere: Concentrations and chemical signatures. *Geophys. Res. Lett.* **2017**, *44* (12), 6226−6234.
(26) Ziolkowski, L. A.; Chamberlin, A. R.; Greaves, J.; Druffel, E. R. M. Quantification of black carbon in marine systems using the benzene polycarboxylic acid method: a mechanistic and yield study. *Limnol. Oceanogr.: Methods* **2011**, *9* (4), 140.
(27) Dittmar, T. The molecular level determination of black carbon in marine dissolved organic matter. *Org. Geochem.* **2008**, *39* (4), 396−407, DOI: 10.1016/j.orggeochem.2008.01.015.
(28) Stubbins, A.; Spencer, R. G. M.; Chen, H.; Hatcher, P. G.; Mopper, K.; Hernes, P. J.; Mwamba, V. L.; Mangangu, A. M.; Wabakanghanzi, J. N.; Six, J. Illuminated darkness: Molecular signatures of Congo River dissolved organic matter and its photochemical alteration as revealed by ultrahigh precision mass spectrometry. *Limnol. Oceanogr.* **2010**, *55* (4), 1467−1477.
(29) He, C.; He, D.; Chen, C.; Shi, Q. Application of Fourier transform ion cyclotron resonance mass spectrometry in molecular characterization of dissolved organic matter. *Sci. China Earth Sci.* **2022**, *65* (12), 2219−2236.
(30) Jaffé, R.; Ding, Y.; Niggemann, J.; Vähätalo, A. V.; Stubbins, A.; Spencer, R. G.; Campbell, J.; Dittmar, T. Global charcoal mobilization from soils via dissolution and riverine transport to the oceans. *Science* **2013**, *340* (6130), 345−347, DOI: 10.1126/science.1231476.
(31) Jones, M. W.; Coppola, A. I.; Santin, C.; Dittmar, T.; Jaffe, R.; Doerr, S. H.; Quine, T. A. Fires prime terrestrial organic carbon for riverine export to the global oceans. *Nat. Commun.* **2020**, *11* (1), No. 2791.
(32) Qi, Y.; Fu, W.; Tian, J.; Luo, C.; Shan, S.; Sun, S.; Ren, P.; Zhang, H.; Liu, J.; Zhang, X.; Wang, X. Dissolved black carbon is not likely a significant refractory organic carbon pool in rivers and oceans. *Nat. Commun.* **2020**, *11* (1), No. 5051.
(33) Olivares, C. I.; Zhang, W.; Uzun, H.; Erdem, C. U.; Majidzadeh, H.; Trettin, C.; Karanfil, T.; Chow, A. Optical in-situ sensors capture dissolved organic carbon (DOC) dynamics after prescribed fire in high-DOC forest watersheds. *Int. J. Wildland Fire* **2019**, *28* (10), 761−768, DOI: 10.1071/WF18175.
(34) Kruger, B. R.; Hausner, M.; Chellman, N.; Weaver, M.; Samburova, V.; Khlystov, A. Dissolved black carbon as a potential driver of surface water heating dynamics in wildfire-impacted regions: A case study from Pyramid Lake, NV, USA. *Sci. Total Environ.* **2023**, *888*, No. 164141.
(35) Chen, Y.; Sun, K.; Sun, H.; Yang, Y.; Li, Y.; Gao, B.; Xing, B. Photodegradation of pyrogenic dissolved organic matter increases bioavailability: Novel insight into bioalteration, microbial community succession, and C and N dynamics. *Chem. Geol.* **2022**, *605*, No. 120964.
(36) Xu, F.; Wei, C.; Zeng, Q.; Li, X.; Alvarez, P. J. J.; Li, Q.; Qu, X.; Zhu, D. Aggregation Behavior of Dissolved Black Carbon: Implications for Vertical Mass Flux and Fractionation in Aquatic Systems. *Environ. Sci. Technol.* **2017**, *51* (23), 13723−13732.
(37) Ou, Q.; Xu, Y.; He, Q.; Wu, Z.; Ma, J.; Huangfu, X. Deposition behavior of dissolved black carbon on representative surfaces: Role of molecular conformation. *J. Environ. Chem. Eng.* **2021**, *9* (5), No. 105921.
(38) Jones, M. W.; de Aragão, L. E. O. C.; Dittmar, T.; Rezende, C. E.; Almeida, M. G.; Johnson, B. T.; Marques, J. S. J.; Niggemann, J.; Rangel, T. P.; Quine, T. A. Environmental Controls on the Riverine Export of Dissolved Black Carbon. *Global Biogeochem. Cycles* **2019**, *33* (7), 849−874, DOI: 10.1029/2018GB006140.
(39) Nakane, M.; Ajioka, T.; Yamashita, Y. Distribution and Sources of Dissolved Black Carbon in Surface Waters of the Chukchi Sea, Bering Sea, and the North Pacific Ocean. *Front. Earth Sci.* **2017**, *5*, No. 34, DOI: 10.3389/feart.2017.00034.
(40) Fang, Z.; Yang, W.; Stubbins, A.; Chen, M.; Li, J.; Jia, R.; Li, Q.; Zhu, J.; Wang, B. Spatial characteristics and removal of dissolved black carbon in the western Arctic Ocean and Bering Sea. *Geochim. Cosmochim. Acta* **2021**, *304*, 178−190.
(41) Ziolkowski, L. A.; Druffel, E. R. M.Aged black carbon identified in marine dissolved organic carbon *Geophys. Res. Lett.* 2010; Vol. *37* 16 DOI: 10.1029/2010GL043963.
(42) Coppola, A. I.; Druffel, E. R. M. Cycling of black carbon in the ocean. *Geophys. Res. Lett.* **2016**, *43* (9), 4477−4482.
(43) Wagner, S.; Cawley, K. M.; Rosario-Ortiz, F. L.; Jaffé, R. In-stream sources and links between particulate and dissolved black carbon following a wildfire. *Biogeochemistry* **2015**, *124* (1−3), 145−161.
(44) Zhao, W.; Bao, H.; Huang, D.; Niggemann, J.; Dittmar, T.; Kao, S.-J. Evidence from molecular marker and FT-ICR-MS analyses for the source and transport of dissolved black carbon under variable water discharge of a subtropical Estuary. *Biogeochemistry* **2023**, *162* (1), 43−55.
(45) Fang, Y.; Chen, Y.; Huang, G.; Hu, L.; Tian, C.; Xie, J.; Lin, J.; Lin, T. Particulate and Dissolved Black Carbon in Coastal China Seas: Spatiotemporal Variations, Dynamics, and Potential Implications. *Environ. Sci. Technol.* **2021**, *55* (1), 788−796.
(46) Fang, Y.; Chen, Y.; Tian, C.; Wang, X.; Lin, T.; Hu, L.; Li, J.; Zhang, G.; Luo, Y. Cycling and Budgets of Organic and Black Carbon in Coastal Bohai Sea, China: Impacts of Natural and Anthropogenic Perturbations. *Global Biogeochem. Cycles* **2018**, *32* (6), 971−986.
(47) Chen, H.; Wang, J.; Zhao, X.; Wang, Y.; Huang, Z.; Gong, T.; Xian, Q. Occurrence of dissolved black carbon in source water and disinfection byproducts formation during chlorination. *J. Hazard. Mater.* **2022**, *435*, No. 129054.
(48) Schneider, E.; Czech, H.; Popovicheva, O.; Lütdke, H.; Schnelle-Kreis, J.; Khodzher, T.; Rüger, C. P.; Zimmermann, R. Molecular Characterization of Water-Soluble Aerosol Particle Extracts by Ultrahigh-Resolution Mass Spectrometry: Observation of Industrial Emissions and an Atmospherically Aged Wildfire Plume at Lake Baikal. *ACS Earth Space Chem.* **2022**, *6* (4), 1095−1107.
(49) Chen, H.; Uzun, H.; Tolić, N.; Chu, R.; Karanfil, T.; Chow, A. T. Molecular Transformation of Dissolved Organic Matter during the Processes of Wildfire, Alum Coagulation, and Disinfection Using

ESI(−) and ESI(+) FT-ICR MS. *ACS ES&T Water* **2023**, *3* (8), 2571−2580.
(50) Wagner, S.; Riedel, T.; Niggemann, J.; Vahatalo, A. V.; Dittmar, T.; Jaffe, R. Linking the Molecular Signature of Heteroatomic Dissolved Organic Matter to Watershed Characteristics in World Rivers. *Environ. Sci. Technol.* **2015**, *49* (23), 13798−13806.
(51) Chen, H.; Wang, J. J.; Ku, P. J.; Tsui, M. T.; Abney, R. B.; Berhe, A. A.; Zhang, Q.; Burton, S. D.; Dahlgren, R. A.; Chow, A. T. Burn Intensity Drives the Alteration of Phenolic Lignin to (Poly) Aromatic Hydrocarbons as Revealed by Pyrolysis Gas Chromatography-Mass Spectrometry (Py-GC/MS). *Environ. Sci. Technol.* **2022**, *56* (17), 12678−12687.
(52) Donahue, N. M.; Epstein, S. A.; Pandis, S. N.; Robinson, A. L. A two-dimensional volatility basis set: 1. organic-aerosol mixing thermodynamics. *Atmos. Chem. Phys.* **2011**, *11* (7), 3303−3318.
(53) Bahureksa, W.; Young, R. B.; McKenna, A. M.; Chen, H.; Thorn, K. A.; Rosario-Ortiz, F. L.; Borch, T. Nitrogen Enrichment during Soil Organic Matter Burning and Molecular Evidence of Maillard Reactions. *Environ. Sci. Technol.* **2022**, *56* (7), 4597−4609.
(54) Peng, Q.; Palm, B. B.; Fredrickson, C. D.; Lee, B. H.; Hall, S. R.; Ullmann, K.; Campos, T.; Weinheimer, A. J.; Apel, E. C.; Flocke, F.; Permar, W.; Hu, L.; Garofalo, L. A.; Pothier, M. A.; Farmer, D. K.; Ku, I. T.; Sullivan, A. P.; Collett, J. L.; Fischer, E.; Thornton, J. A. Observations and Modeling of NOx Photochemistry and Fate in Fresh Wildfire Plumes. *ACS Earth Space Chem.* **2021**, *5* (10), 2652−2667.
(55) Lindaas, J.; Pollack, I. B.; Garofalo, L. A.; Pothier, M. A.; Farmer, D. K.; Kreidenweis, S. M.; Campos, T. L.; Flocke, F.; Weinheimer, A. J.; Montzka, D. D.; Tyndall, G. S.; Palm, B. B.; Peng, Q.; Thornton, J. A.; Permar, W.; Wielgasz, C.; Hu, L.; Ottmar, R. D.; Restaino, J. C.; Hudak, A. T.; Ku, I. T.; Zhou, Y.; Sive, B. C.; Sullivan, A.; Collett, J. L.; Fischer, E. V. Emissions of Reactive Nitrogen From Western U.S. Wildfires During Summer 2018. *J. Geophys. Res.: Atmos.* **2021**, *126* (2), No. e2020JD032657.
(56) Schneider, E.; Czech, H.; Popovicheva, O.; Chichaeva, M.; Kobelev, V.; Kasimov, N.; Minkina, T.; Rüger, C. P.; Zimmermann, R. Comprehensive mass spectrometric analysis of unprecedented high levels of carbonaceous aerosol particles long-range transported from wildfires in the Siberian Arctic *EGUsphere* 2023 DOI: 10.5194/egusphere-2023-769.
(57) Song, J.; Li, M.; Fan, X.; Zou, C.; Zhu, M.; Jiang, B.; Yu, Z.; Jia, W.; Liao, Y.; Peng, P. Molecular Characterization of Water- and Methanol-Soluble Organic Compounds Emitted from Residential Coal Combustion Using Ultrahigh-Resolution Electrospray Ionization Fourier Transform Ion Cyclotron Resonance Mass Spectrometry. *Environ. Sci. Technol.* **2019**, *53* (23), 13607−13617.
(58) Laskin, A.; Smith, J. S.; Laskin, J. Molecular characterization of nitrogen-containing organic compounds in biomass burning aerosols using high-resolution mass spectrometry. *Environ. Sci. Technol.* **2009**, *43* (10), 3764−3771.
(59) Bateman, A. P.; Nizkorodov, S. A.; Laskin, J.; Laskin, A. High-resolution electrospray ionization mass spectrometry analysis of water-soluble organic aerosols collected with a particle into liquid sampler. *Anal. Chem.* **2010**, *82* (19), 8010−8016.
(60) Wagner, S.; Dittmar, T.; Jaffé, R. Molecular characterization of dissolved black nitrogen via electrospray ionization Fourier transform ion cyclotron resonance mass spectrometry. *Org. Geochem.* **2015**, *79*, 21−30.
(61) Ditto, J. C.; He, M.; Hass-Mitchell, T. N.; Moussa, S. G.; Hayden, K.; Li, S.-M.; Liggio, J.; Leithead, A.; Lee, P.; Wheeler, M. J.; Wentzell, J. J. B.; Gentner, D. R. Atmospheric evolution of emissions from a boreal forest fire: the formation of highly functionalized oxygen-, nitrogen-, and sulfur-containing organic compounds. *Atmos. Chem. Phys.* **2021**, *21* (1), 255−267.
(62) Gomez-Saez, G. V.; Pohlabeln, A. M.; Stubbins, A.; Marsay, C. M.; Dittmar, T. Photochemical Alteration of Dissolved Organic Sulfur from Sulfidic Porewater. *Environ. Sci. Technol.* **2017**, *51* (24), 14144−14154.
(63) Stubbins, A.; Niggemann, J.; Dittmar, T. Photo-lability of deep ocean dissolved black carbon. *Biogeosciences* **2012**, *9* (5), 1661−1670.
(64) Liu, M.; Zhao, Z.; Lu, Q.; Yu, W. Release of dissolved organic carbon from biochar and formation of humic-like component during photoreaction: Effects of Ca(2+) and pH. *Water Res.* **2022**, *219*, No. 118616.
(65) Hutchison, K. J.; Hesterberg, D.; Chou, J. W. Stability of Reduced Organic Sulfur in Humic Acid as Affected by Aeration and pH. *Soil Sci. Soc. Am. J.* **2001**, *65* (3), 704−709.
(66) Poulin, B. A. Selective Photochemical Oxidation of Reduced Dissolved Organic Sulfur to Inorganic Sulfate. *Environ. Sci. Technol. Lett.* **2023**, *10* (6), 499−505.
(67) Hesterberg, D.; Chou, J. W.; Hutchison, K. J.; Sayers, D. E. Bonding of Hg (II) to reduced organic sulfur in humic acid as affected by S/Hg ratio. *Environ. Sci. Technol.* **2001**, *35* (13), 2741−2745.
(68) Zhou, Z.; Chen, B.; Qu, X.; Fu, H.; Zhu, D. Dissolved Black Carbon as an Efficient Sensitizer in the Photochemical Transformation of 17beta-Estradiol in Aqueous Solution. *Environ. Sci. Technol.* **2018**, *52* (18), 10391−10399.
(69) Yan, W.; Chen, Y.; Han, L.; Sun, K.; Song, F.; Yang, Y.; Sun, H. Pyrogenic dissolved organic matter produced at higher temperature is more photoactive: Insight into molecular changes and reactive oxygen species generation. *J. Hazard. Mater.* **2022**, *425*, No. 127817.
(70) Moran, M. A.; Durham, B. P. Sulfur metabolites in the pelagic ocean. *Nat. Rev. Microbiol.* **2019**, *17* (11), 665−678.
(71) Soleimani, M.; Bassi, A.; Margaritis, A. Biodesulfurization of refractory organic sulfur compounds in fossil fuels. *Biotechnol. Adv.* **2007**, *25* (6), 570−596.
(72) Berman, T.; Bronk, D. A. Dissolved organic nitrogen: a dynamic participant in aquatic ecosystems. *Aquat. Microb. Ecol.* **2003**, *31* (3), 279−305.

https://doi.org/10.1021/acs.est.4c00492
Environ. Sci. Technol. XXXX, XXX, XXX−XXX